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Journal Name	Boundary-Layer Meteorology	
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Schedule	Received	28 October 2010
	Revised	
	Accepted	21 February 2012

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## Abstract

Aerosol dispersion in the area surrounding an existing biological treatment facility is investigated using large-eddy simulation, with the objective to investigate the applicability of computational fluid dynamics to complex real-life problems. The aerosol sources consist of two large aeration ponds that slowly diffuse aerosols into the atmosphere. These sources are modelled as dilute concentrations of a non-buoyant non-reacting pollutant diffusing from two horizontal surfaces. The time frame of the aerosol release is restricted to the order of minutes, justifying a statistically steady inlet boundary condition. The numerical results are compared to wind-tunnel experiments for validation. The wind-tunnel flow characteristics resemble neutral atmospheric conditions with a Reynolds number, based on the boundary-layer thickness, of  $Re_{\delta} \approx 2 \times 10^5$ . The numerical inflow conditions are based upon the wind-tunnel flow field. The predicted decay of both the mean and root-mean-square concentrations are in good agreement with experimental data; at 3 m from the ground, the plume mean concentration 200 m downwind of the source is approximately 2% of the source strength. The numerical data in the near-surface layer (0–50 m from the ground) correspond particularly well with the wind-tunnel data. Tentative deposition simulations suggest that there seems to be little difference in the deposition rates of large ( $1.8 \times 10^{-5}$  m) and small ( $3 \times 10^{-6}$  m) particles in the near-field under the flow conditions considered.

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Keywords (separated by '-') Aerosols - Dispersion - Large-eddy simulation - Turbulence - Validation

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Footnote Information

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# On the Use of Computational Fluid Dynamics to Investigate Aerosol Dispersion in an Industrial Environment: A Case Study

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Received: 28 October 2010 / Accepted: 21 February 2012  
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**Abstract** Aerosol dispersion in the area surrounding an existing biological treatment facility is investigated using large-eddy simulation, with the objective to investigate the applicability of computational fluid dynamics to complex real-life problems. The aerosol sources consist of two large aeration ponds that slowly diffuse aerosols into the atmosphere. These sources are modelled as dilute concentrations of a non-buoyant non-reacting pollutant diffusing from two horizontal surfaces. The time frame of the aerosol release is restricted to the order of minutes, justifying a statistically steady inlet boundary condition. The numerical results are compared to wind-tunnel experiments for validation. The wind-tunnel flow characteristics resemble neutral atmospheric conditions with a Reynolds number, based on the boundary-layer thickness, of  $Re_\delta \approx 2 \times 10^5$ . The numerical inflow conditions are based upon the wind-tunnel flow field. The predicted decay of both the mean and root-mean-square concentrations are in good agreement with experimental data; at 3 m from the ground, the plume mean concentration 200 m downwind of the source is approximately 2% of the source strength. The numerical data in the near-surface layer (0–50 m from the ground) correspond particularly well with the wind-tunnel data. Tentative deposition simulations suggest that there seems to be little difference in the deposition rates of large ( $1.8 \times 10^{-5}$  m) and small ( $3 \times 10^{-6}$  m) particles in the near-field under the flow conditions considered.

**Keywords** Aerosols · Dispersion · Large-eddy simulation · Turbulence · Validation

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## 1 Introduction

Borregaard Ind. Ltd., located in Sarpsborg, Norway, is the world's leading supplier of wood-based chemicals, and waste water from their wood refinement plant is biologically treated according to environmental requirements legislated by the Norwegian Environmental Protection Agency. Here we consider the continuous release of aerosols from such a biological treatment facility.

The facility consists of two large open-air aeration ponds each containing about 30,000 m<sup>3</sup> of liquid, each having a diameter of approximately 40 m. In order to promote optimal growth of microorganisms achieving efficient degradation of various organic substances, the pond temperature is maintained constant at about 37°C. In addition, about 8.3 m<sup>3</sup> s<sup>-1</sup> of constant temperature air (≈20°C) is circulated through each pond, giving rise to massive aerosol generation at the liquid surface. These aerosols, which potentially may contain microorganisms from the ponds, are emitted with low vertical velocity into the surroundings, and then transported and dispersed downwind by the wind field.

Liquid samples taken directly from these aeration ponds (Blatny et al. 2008, 2011), and samples of aerosols collected immediately above the liquid surfaces (Olsen et al. 2010), have all revealed the presence of the microorganism *Legionella* spp. Blatny et al. (2008) proved that *Legionella*-containing aerosols could be dispersed from the aeration ponds and transported at least 300 m downwind of the ponds. Epidemiological investigations in conjunction with outbreaks of Legionnaires disease among the population (Nguyen et al. 2006; Nygård et al. 2008) have even suggested that *Legionella*-containing aerosols may be transported up to 10 km from its source. The true source of *Legionella*-containing airborne aerosols that caused the largest outbreak of Legionnaires disease in Norway by inhalation has, however, not yet been unambiguously identified. During an outbreak in May, 2005, 10 people died and 56 patient cases were documented (Nygård et al. 2008).

The present study constitutes an integral part of a broader initiative to gain further insight into the dispersion and transport processes, and the microbiological characteristics, of aerosols emitted from the aeration ponds at Borregaard Ind. Ltd., Norway (cf. Blatny et al. 2008, 2011; Olsen et al. 2010). The aeration ponds considered in this study have been identified as two of several possible sources responsible for the outbreak in 2005.

The *Legionella* outbreaks occurred over several days. During the 2005 outbreak, the meteorological data generally showed wind conditions corresponding to a Reynolds number, based on the boundary-layer thickness  $\delta$ , of  $Re_\delta \sim 10^8$  and neutral to unstable conditions. The flow characteristics used herein were chosen such that the wind-tunnel experiments and computer simulations ensure that conditions are no more favourable to aerosol mixing than the actual meteorological conditions during the outbreak. This was achieved by considering isothermal conditions at a similar  $Re_\delta$ .

There exists a plethora of modelling approaches to simulate aerosol dispersion. Simple computational models based on e.g. the Gaussian-puff approach (Sykes and Gabruk 1996) are frequently used to characterize contaminant dispersion in the atmospheric boundary layer. These should normally not be applied in microscale complex urban or industrial environments since the fine-scale geometrical features are of crucial importance.

Fuelled by the need for «fast» urban dispersion models, a number of approaches to extrapolate the use of the puff models to urban terrain have been proposed (e.g. Taylor and Salmon 1993; Kaplan and Dinar 1996). Although progress has been made in recent years, there are still very large uncertainties associated with the use of these techniques in complex built-up terrain. Despite its computational cost, computational fluid dynamics (CFD) is increasingly being used in urban dispersion applications. As pointed out by Lee et al. (2000), however,

67 the application of CFD in urban environments highlights a number of modelling issues that  
68 need to be addressed in order to warrant its use. Many CFD studies addressing dispersion in  
69 urban or industrial environments are based on the assumption that the flow field is statisti-  
70 cally steady. The steady-state Reynolds-averaged Navier–Stokes (RANS) approach is thus  
71 frequently used, cf. e.g. Lien and Yee (2004), Coirier et al. (2005), Lien et al. (2006), Santiago  
72 et al. (2007).

73 Lien et al. (2006) concluded that the standard linear eddy-viscosity  $k-\varepsilon$  model was perhaps  
74 the simplest complete turbulence model that could be used for urban flow predictions. They  
75 also noted, however, that while the mean velocity field can be quite well predicted using  
76 this approach, the turbulence kinetic energy seems in general to be under-predicted. It thus  
77 follows that also the mixing process caused by irregular velocity fluctuations may be poorly  
78 represented. Ad hoc modifications of the RANS model can be devised to improve the pre-  
79 dictions of the turbulence kinetic energy, cf. e.g. Wang et al. (2008), though the general  
80 applicability of these modifications to arbitrary complex building structures and variation in  
81 topography are questionable.

82 As already alluded to, the application of CFD for dispersion and transport modelling in  
83 complex urban or industrial environments are confronted with many challenges. The effects  
84 on the flow field of the interaction between the blockage caused by building structures and  
85 the atmospheric boundary layer (ABL) are crucially important and need to be accounted for  
86 in the computational model.

87 From a modelling perspective, the dominating effects are, (i) kinematic blocking of veloc-  
88 ity components normal to solid surfaces, and (ii) non-local effects caused by pressure reflec-  
89 tions (cf. e.g. Durbin and Reif 2002). The latter dominates in the ABL where it greatly  
90 modifies the turbulence anisotropy and consequently also the dispersion process. Kinematic  
91 blocking dominates the local flow conditions in densely built-up areas with tall buildings and  
92 produces street canyon (channeling) effects, flow separation, and the subsequent generation  
93 of unsteady wakes.

94 The mesoscale atmospheric events affecting the incoming ABL flow can, to a good approx-  
95 imation, be treated as statistically steady when considering releases of up to an hour. The  
96 dominating fluid dynamical processes within urban terrain are however essentially statisti-  
97 cally unsteady.

98 There are three dominating sources of this flow unsteadiness: (i) instantaneous spatial and  
99 temporal variations of the energy containing large-scale motion of the incoming flow, (ii) the  
100 flow in the wake region downstream of buildings, and (iii) self-generated unsteadiness related  
101 to the emission itself. The ABL is characterized by a continuous range of spatial scales from  
102 a few tens of millimetres up to several hundreds of metres, with temporal scales varying  
103 from a fraction of a second up to a few tens of minutes. The large-scale time variation of  
104 building wakes can be up to an order of some minutes. These temporal effects are crucially  
105 important in cases where the source release time is in the order of, say, one hour or less,  
106 which is typically the case for accidental or intentional releases.

107 In an effort to overcome the challenges associated with unsteady mixing, which is believed  
108 to play a crucial role in order to accurately predict contaminant transport, the present study  
109 adopts the large-eddy simulation (LES) approach that, at least conceptually, resolves the tem-  
110 poral and spatial variation of the large energetic scales in the turbulent boundary layer. The  
111 objective is to increase our understanding of the near-field transport (<500 m) of aerosols  
112 emitted from aeration ponds in general, and for those at Borregaard Ind. Ltd. specifically, that  
113 potentially carry microbiological material that may result in disease. The present study thus  
114 addresses a real-life scenario characterized by a considerable separation of spatial and tem-  
115 poral scales. The computational effort uses what is practically achievable with the available

116 computer resources. The simulation results are verified by data from recent wind-tunnel tests  
 117 carried out at the Environmental Flow Research Laboratory (EnFlo) at the University of  
 118 Surrey, Guildford, UK.

119 The primary objectives are (i) to verify whether CFD is a suitable tool for inves-  
 120 tigating near-field aerosol transport in the ABL, and if so (ii) to use CFD results to  
 121 shed light on the role of aeration ponds during the largest outbreak of Legionella ever  
 122 recorded in Norway. The conclusions drawn will be determined by how well the com-  
 123 putational model predicts various aspects of the flow and concentration fields. As such,  
 124 comments regarding the physical phenomena of the flow will also be included in the  
 125 analysis of the results, as well as in the conclusions. In the present study these phe-  
 126 nomena are discussed in the context of model verification, not as an independent topic.  
 127 However, if a computational model is to be of any use in later investigations into the  
 128 physics of aerosol dispersion, it is imperative that realistic physical phenomena emerges  
 129 from the model. Hence, part of this paper also focuses on the dispersion features them-  
 130 selves.

## 131 2 Problem Characterization

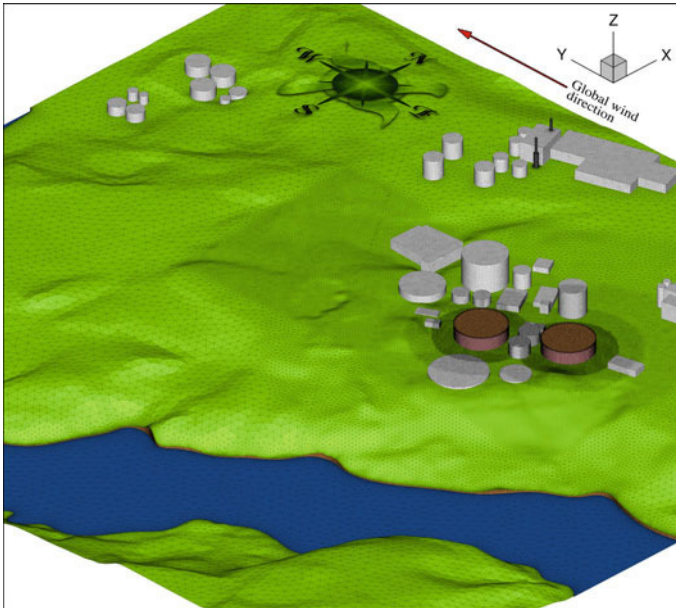
### 132 2.1 Geometry

133 The full-scale dimensions of the computational domain ( $L_x \times L_y \times L_z$ ) are approximately  
 134  $700 \times 800 \times 400 \text{ m}^3$ , where the global wind direction is in the  $y$ -direction and  $z$  is the vertical  
 135 direction. The wind-tunnel model is of scale 1:300, i.e. approximately  $2.3 \times 2.7 \times 1.3 \text{ m}^3$ , and  
 136 these are also the actual dimensions used in the computer simulation model. Most buildings  
 137 in the area belong to the industrial plant. The surrounding landscape consists of grasslands  
 138 with a few trees, some sloping hills and a river, see Fig. 1. See also Sect. 2.3 for topographical  
 139 details, and Sect. 3.1 for details of the computational mesh.

140 Aerosols are released from the two aeration ponds marked in red (dark) in Fig. 1. The  
 141 full-scale height of the ponds is approximately 12 m, with diameter around 42 m. As thermal  
 142 effects are ignored and the source is considered dilute, no buoyancy or two-way coupling  
 143 effects are likely to be present. In real life, the ponds contain water ( $\approx 37^\circ\text{C}$ ), from which  
 144 aerosols diffuse into the atmosphere with very small vertical velocity. It is assumed that  
 145 the temperature difference between the pond and the surrounding air has an insignificant  
 146 impact on the surrounding flow. The aerosols are considered small and their concentration  
 147 sufficiently dilute that they do not influence the wind field.

### 148 2.2 The Atmospheric Boundary Layer

149 As previously stated, the current numerical simulation and wind-tunnel experiment are per-  
 150 formed under neutral conditions, assuming constant temperature in the entire domain. Potent-  
 151 tial humidity effects are also neglected. These conditions ensure at least as much mixing as  
 152 the real-life conditions observed during the outbreaks, in particular the largest in 2005. It  
 153 should be recalled that the meteorological conditions during the outbreak were neutral to  
 154 unstable. The main motivation for using such conditions in the wind tunnel is to derive a  
 155 conservative estimate for the maximum mean concentration of Legionella downstream of  
 156 the source. The incoming wind field is created in the wind tunnel so as to imitate a real-life  
 157 atmospheric field (cf. e.g. [Robins et al. 2001](#)) with a Reynolds number  $Re_\delta \approx 2 \times 10^5$ . In



**Fig. 1** The computational mesh of the Borregaard Ind. Ltd. geometry. Full-scale dimensions ( $L_x \times L_y \times L_z$ ) are approximately  $700 \times 800 \times 400 \text{ m}^3$ . Cell sizes range from 0.5 to 7.5 m. The aeration ponds are marked in red

158 the computer simulation, the incoming wind field is generated using wind-tunnel data (see  
159 Sect. 3.4.1).

160 Exact data for the atmospheric conditions during the Leigionella outbreak in 2005 are not  
161 available. However, during the entire period of the outbreaks (about a week), the wind field  
162 was in the high Reynolds number range ( $Re_\delta \sim 10^8$ ), and the boundary layer was neutral  
163 or unstable. The Reynolds number in the wind tunnel is thus lower than that of the real-life  
164 atmospheric boundary layer, though the flow is still in the high Reynolds number range,  
165 in which scalar dispersion is relatively independent of the Reynolds number. As mentioned  
166 previously, in the context of the present study, an approach that ensures the same or greater  
167 mixing than the real case is more important than merely attempting to recreate or model  
168 real-life conditions exactly.

169 For validation purposes the same conditions as in the wind tunnel must be used in the  
170 computer simulation. Once successfully verified, the computational approach may be used  
171 to conduct further studies in which e.g. the atmospheric conditions are altered, or different  
172 aerosol sizes are considered. This is, however, outside the scope of the present study.

### 173 2.3 Topography

174 The present topography differs significantly both from entirely unpopulated areas and densely  
175 populated cities. However, some conclusions drawn from this study will potentially have  
176 implications for similar problems with domains of comparable sizes. Indeed, the verifica-  
177 tion of simulation methodology obtained herein will in large part be valid also for different  
178 topographies. Note that the results themselves, though, cannot be translated to different



**Table 1** Key features of the topography, in full scale

Building count	45
Average building area	539.43 m <sup>2</sup>
Average building height	10.80 m
Standard deviation of building height	6.57 m
Building density for bounding box of buildings	1.76 × 10 <sup>-4</sup> bld m <sup>-2</sup>
Building density for entire domain	8.19 × 10 <sup>-5</sup> bld m <sup>-2</sup>
Minimum landscape elevation	-3 m
Maximum landscape elevation	53 m
Average landscape elevation	21.9 m
Standard deviation of landscape elevation	17.6 m

179 topographies. A few basic features of the topography used in this study are nevertheless  
180 summarized in Table 1.

### 181 3 Computational Modelling

#### 182 3.1 Mesh Generation

183 A mesh consisting of  $4.2 \times 10^7$  cells has been used to discretize the geometry, cf. Fig. 1.  
184 The height of the domain is the same as in the wind tunnel, but the computational domain  
185 is smaller than the wind-tunnel dimensions in the horizontal directions. Adequate space  
186 between the region of interest and the domain boundaries is ensured, following the directions  
187 of the COST Action 732 guidelines (cf. e.g. Franke et al. 2011).

188 Comparisons with simulations on a coarser mesh<sup>1</sup> suggested that the flow patterns glob-  
189 ally were reasonably unaffected by the grid. It should be noted that this case constitutes an  
190 approximation of a real-life scenario with significant scale separations. Grid independence  
191 is virtually impossible, but the assessment of the present grid suggests that it resolves the  
192 most dominant flow features. Reports of non-dimensional distance  $y_+ = d_y u_* / \nu$  from the  
193 ground, where  $u_* = (\tau_w / \rho)^{1/2}$  is the friction velocity given in terms of density  $\rho$  and wall  
194 stress  $\tau_w$ ,  $d_y$  is distance from the wall, and  $\nu$  is kinematic viscosity, implied that the near-wall  
195 resolution was adequate ( $y_+ < 50$ ) for a built environment.

196 The aerosol sources, i.e. the surfaces of the aeration ponds, each comprised more than  
197 13,000 cell faces, with a surrounding grid resolution of approximately 0.5 m.

#### 198 3.2 Turbulence Modelling

The LES model, utilizing the dynamic Smagorinsky approach, has been used, as imple-  
mented in the commercial software ANSYS Fluent 12.0 (Fluent 6.3 User's Guide 2006). For  
a Newtonian, incompressible flow, the filtered equations for conservation of momentum and  
mass can be written as

$$\frac{\partial U_i}{\partial t} + U_j \frac{\partial U_i}{\partial x_j} = -\frac{1}{\rho} \frac{\partial P}{\partial x_i} + \nu \frac{\partial^2 U_i}{\partial x_j \partial x_j} - \frac{\partial \tau_{ij}}{\partial x_j}, \quad (1)$$

$$\frac{\partial U_i}{\partial x_i} = 0, \quad (2)$$

<sup>1</sup> The overall cell density in the coarse mesh was only 14% of that in the fine mesh. However, near the ground, the coarse mesh had a cell density of about 70% of the fine mesh.

199 respectively, where  $U_i(\mathbf{x}, t)$  and  $P(\mathbf{x}, t)$  denote the filtered velocity component in the  $x_i$   
 200 direction and the filtered pressure field, respectively, and  $\nu$  is the kinematic viscosity.  $\tau_{ij} =$   
 201  $u'_i u'_j - U_i U_j$  is the residual stress, or subgrid-scale, tensor, which physically represents the  
 202 impact of the unresolved velocity components on the computationally resolved velocity field.  
 203 The present study utilizes the Smagorinsky model (cf. e.g. Kim 2004), given by

$$204 \quad \tau_{ij} - \frac{1}{3} \tau_{kk} \delta_{ij} = -2\nu_t \bar{S}_{ij}, \quad (3)$$

205 where  $\nu_t = L_s^2 |S|$  is the subgrid-scale turbulent viscosity. The length scale is  $L_s =$   
 206  $\min(\kappa d_y, C_s V^{1/3})$ , where  $\kappa$  is the von Kármán constant,  $d_y$  is the distance to the closest  
 207 wall, and  $V$  is the volume of the computational cell. The Smagorinsky constant  $C_s$  is deter-  
 208 mined dynamically and is computed based on the resolved scales of motion. The rate-of-strain  
 209 tensor for the resolved scale is given by  $\bar{S}_{ij} = \frac{1}{2} (\frac{\partial U_i}{\partial x_j} + \frac{\partial U_j}{\partial x_i})$ .

### 210 3.3 Dispersion Modelling

211 Two different dispersion modelling strategies have been adopted. Firstly, a scalar-transport  
 212 model has been employed to simulate the plume evolutions. Secondly, a Lagrangian discrete-  
 213 particle approach has been used to investigate particle ground deposition. Experimental data  
 214 are used to validate the plume evolution up to distances of 500m downwind of the aeration  
 215 ponds (see Sect. 4), and representative aerosol sizes experimentally obtained from full-scale  
 216 measurements (Blatny et al. 2011) are used in the Lagrangian approach.

#### 217 3.3.1 Scalar Transport

218 The transport of passive scalars, such as the tracer gas used in the wind tunnel, is governed  
 219 by the advection-diffusion equation

$$220 \quad \frac{\partial c}{\partial t} + U_i \frac{\partial c}{\partial x_i} = \nabla^2 (\alpha_c c + \frac{\nu_T}{Sc_T} c), \quad (4)$$

221 where  $c = c(\mathbf{x}, t)$  represent the scalar field and  $\alpha_c$  denotes the scalar molecular diffusivity,  
 222 and the last term represents a model for subgrid-scale turbulent diffusion. In the present study,  
 223 the subgrid-scale turbulent Schmidt number  $Sc_T$  is obtained through the dynamic procedure  
 224 described in Germano et al. (1991).

#### 225 3.3.2 Particle Transport

226 The discrete-particle transport model incorporates inertia and gravity effects inherently  
 227 ignored in the scalar transport model. The displacement of a given particle follows from  
 228 integrating the force balance on that particle, which can be written in a Lagrangian reference  
 229 frame. For example, in the tangential direction  $x_p$  of the particle trajectory, the force balance  
 230 can be written as

$$231 \quad \frac{du_p}{dt} = F_D(u_{x_p} - u_p) + \frac{g_{x_p}(\rho_p - \rho)}{\rho_p} + F_{x_p}, \quad (5)$$

232 where  $u_{x_p}$  and  $u_p$  are the fluid and particle velocities, respectively, in the  $x_p$  direction. The  
 233 former is obtained from the LES flow field.  $F_D$  denotes the drag force in the  $x_p$  direction

on the particle exerted by the wind field,  $\rho$  and  $\rho_p$  are the densities of the fluid and the particle, respectively, and  $g_{x_p}$  is the gravitational acceleration in the  $x_p$  direction.  $F_{x_p}$  involves pressure-gradient and virtual-mass effects, i.e.

$$F_{x_p} = \frac{1}{2} \frac{\rho}{\rho_p} \frac{d}{dt} (u_{x_p} - u_p) + \left( \frac{\rho}{\rho_p} \right) u_{p,i} \frac{\partial u_{x_p}}{\partial x_{p,i}} \quad (6)$$

The drag force is modelled using a spherical drag law, i.e.

$$F_D = \frac{18\mu}{\rho_p d_p^2} \frac{C_D Re}{24}, \quad (7)$$

where  $d_p$  is the diameter of the particle, and  $C_D$  is the drag coefficient, in the present study given by  $C_D = a_1 + a_2/Re + a_3/Re^2$ . The constants  $a_1$ ,  $a_2$  and  $a_3$  apply to smooth spherical particles over several spatial ranges given by Morsi and Alexander (1972). Also in the above,  $Re$  is the slip velocity Reynolds number, defined as  $Re \equiv \rho d_p |u_p - u_{x_p}| / \mu$ .

In its original form, integration of Eq. 5 only yields the particle velocity based on the resolved scales of motion. To account for the subgrid scales, a stochastic model is applied, through which a fluctuating velocity component is added to the fluid velocity used in Eq. 5 at every timestep. The fluctuating velocity components are discrete piecewise constant functions of time, where their random value is kept constant over an interval of time given by the characteristic lifetime of the eddies, computed from the resolved field.

In this study, thermophoretic forces, Brownian forces, and lift forces are assumed to be insignificant and are therefore neglected. Also, effects such as evaporation, agglomeration, and particle collisions are neglected. The particle release in this study can be considered dilute, making these simplifying assumptions approximately valid (Elgobashi 1994).

### 3.4 Numerical Approach

The pressure-based solver in Fluent is used. The second-order accurate bounded central-differencing scheme has been used for the momentum equations, while the pressure-velocity coupling is handled by the semi-implicit method for pressure-linked equations (commonly known as the SIMPLE scheme). A second-order implicit Euler temporal scheme is used. The Lagrangian transport equations are integrated through a combination of a lower order implicit Euler scheme, which is unconditionally stable, and a higher-order semi-implicit trapezoidal scheme. An automated switch determines which scheme to use, depending on how far the particle is from hydrodynamic equilibrium.

The temporal resolution,  $\Delta t$ , was estimated using the large-scale turbulent time scale, i.e.  $\tau = L/U = 0.0146$  s, where  $U$  is the mean inlet velocity at height  $L$  from the ground. The timestep was set to be  $\Delta t = \tau$ . The turbulent length scale  $L$ , also used to determine eddy sizes at the inlet, is estimated as 2/3 of pond height, i.e. approximately 7.3 m (full scale), cf. Fig. 1.

#### 3.4.1 Boundary Conditions

All vertical boundaries of the domain, except building walls, are set as pressure outlets, which allows for an unsteady flow evolution, and they allow particles to escape. No-slip conditions are used on solid surfaces, such as the ground, and walls and roofs of the buildings. The top surface of the computational domain is defined as a symmetry plane, with an escape condition for discrete particles. The top boundary is sufficiently far from the ground, such that any influence of the symmetry assumption is negligible on the solution. The pond surfaces,

where the aerosol generation takes place, are set as no-slip walls, but with a concentration of 100% polluted air (with the same dynamical properties as air), and a reflection condition for discrete particles that are placed on the surfaces when running the discrete-particle model.

The inlet conditions required special consideration in order to be comparable to the experimental conditions. The measured wind profile at a point close to the inlet of the simulation domain was interpolated by a cubic function up to the maximum measuring height. Above this height, the free-stream velocity from the wind-tunnel inlet was used.

The turbulence properties at the inlet were estimated via measurements of velocity-fluctuation correlations in the streamwise and vertical directions, taken at the same position as the mean flow-field measurements. That is,  $\overline{v'^2}$ ,  $\overline{w'^2}$  and  $\overline{v'w'}$  were set as the averages of measured experimental values, where  $v'$  is the streamwise component and  $w'$  is the vertical component. The remaining subgrid stress components were determined by assuming axisymmetric turbulence, i.e.  $\overline{u'^2} = \overline{w'^2}$ ,  $\overline{u'v'} = \overline{v'w'}$  and  $\overline{u'w'} = 0$ . It should be noted that due to lack of experimental data, the present velocity inlet conditions are only approximate. The presence of buildings, however, is believed to reduce the effect of the inlet profile, at least close to the ground.

Finally, from comparisons of streamwise velocity fluctuations  $v'$  to the mean free-stream streamwise velocity  $V$  at the measuring location used for the inlet condition, it was found that  $|v'/V| \approx 0.1$ , implying a turbulence intensity of about 10%. This, along with the aforementioned length scale  $L$ , was also specified as a parameter used to generate eddies at the inlet through a spectral synthesizer method.

### 3.4.2 Statistics

Prior to the statistical sampling, the flow evolved for 4,000 timesteps (i.e. for 1 min) from its initial conditions. Statistics were gathered for more than 14,000 timesteps, representing approximately 1,400 uncorrelated samples, or 3.5 min of simulation time. It is important to note that only the statistics of the flow and dispersion are considered in this study. There is virtually an infinite number of possible realizations of the flow pattern in real life, hence only a statistical prediction is meaningful. In that respect, care must be taken to collect statistics long enough to obtain converged statistical data, but rapid enough that boundary conditions do not change significantly. In the case of a controlled computer simulation and wind-tunnel experiment, as is the case in the present study, such conditions can be controlled and thus cause no problems.

For the discrete particle transport, preliminary tests showed that the same timestep could be used as for the flow simulations. During an initial phase of five timesteps (0.073 s), about 275,000 particles were placed at rest on the surfaces of the two aeration ponds. The total number of released particles in the simulation was equally divided between two dominating particle sizes obtained in recent field measurements (cf. Blatny et al. 2011): 18 and 2  $\mu\text{m}$ . After the initial phase, all particles were tracked until most (>99%) had left the computational domain or were trapped.

It should be noted that each particle in the simulation does not necessarily correspond to one particle in full scale. In fact, a scaling with the measured particle count, preferably at several locations, is necessary to provide the real particle counts (e.g. in parts per million or some other measure). However, regardless of scaling, once the simulation method has been verified, the simulations provide valuable information about the pattern and speed of the propagation of particles.

In order to compare the computational results with wind-tunnel concentration measurements it was necessary to use the scalar field approach. The statistical convergence using particles is very slow and thus computationally demanding, and the scalar transport approach offers a viable alternative. These two approaches are in fact formally equivalent in the limit of statistical convergence if particle mass can be neglected (and particle collisions, agglomeration and break-ups ignored). The latter requirement is fulfilled here since  $St \ll 1$ . The scalar field approach can, however, not be used to quantify aerosol deposition. The discrete particle approach was therefore adopted.

It should finally be noted that also in the case of species transport,  $\Delta t$  was set equal to the flow simulation timestep, again verified by similar methods as in the flow calculations. Statistics were gathered for more than 14,000 time steps (about 3.5 min).

#### 4 Wind-Tunnel Measurements

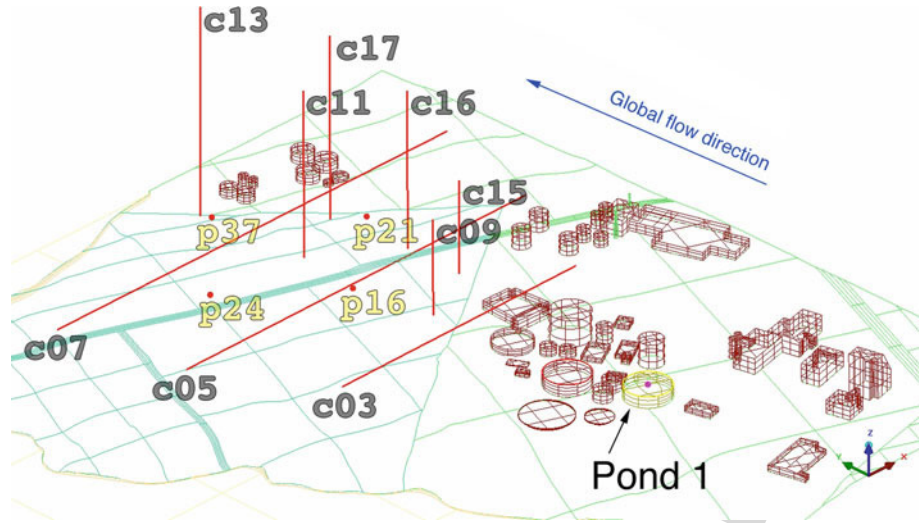
A 1:300 scale wind-tunnel model of the terrain and buildings surrounding Borregaard was manufactured using high density polyurethane foam. The model, which represents approximately  $1 \text{ km}^2$  in full scale, was assembled in the large atmospheric wind tunnel at the Environmental Flow Research Centre at the University of Surrey, Guildford, UK (cf. e.g. Robins et al. 2001). The wind tunnel has a 20 m long test section with a cross-sectional area of  $3.5 \times 1.5 \text{ m}^2$ , with the first 10 m used for flow conditioning purposes. The tunnel can be operated at a free-stream velocity between 0.8 and  $2.5 \text{ m s}^{-1}$ . A large number of planar roughness elements was mounted on the wind-tunnel floor upstream of the model to create a realistic incoming boundary layer. The circular model (3.4 m in diameter) was mounted on a rotatable disk on the tunnel floor whereby different flow directions could be tested by rotating the model.

In this study we focus on one flow direction and one flow speed (corresponding to  $Re_\delta = 2 \times 10^5$ ). The wind-tunnel model incorporates many details of the terrain, including the river Glomma, and the major building structures at Borregaard, but for practical purposes the model had to be simplified. For instance forested areas and other vegetation present on site or in its immediate vicinity were removed, as well as small buildings. The wind-tunnel measurements presented here are limited to isothermal atmospheric conditions.

The released contaminant consisted of a mixture of air with a trace of propane ( $\text{C}_3\text{H}_8$ ). In order to simulate the aerosol generation at the surface of the ponds, each pond was filled with a large number of small spherical high-density polyethylene balls (3 mm in diameter) in order to create a porous media through which the tracer gas could slowly diffuse into the surroundings. The tracer could thus be released from the aeration ponds uniformly over the cross-sectional area of each pond and with a very small vertical velocity.

The choice of using a tracer gas in place of particles is justified in this case since the particle inertia can be neglected. The Stokes number  $St$ , defined as the ratio between the particle relaxation time ( $\tau = \rho_p d_p^2 / (18\mu)$ ) and the characteristic time scale ( $t$ ) for the wind field, is  $\ll 1$ . The particle relaxation time scale is computed using the aerosol ( $\text{H}_2\text{O}$ ) density ( $\rho_p = 998 \text{ kg m}^{-3}$  at  $20^\circ\text{C}$ ), the molecular viscosity of the surrounding air ( $\mu_{\text{air}} = 1.67 \times 10^{-5} \text{ N s m}^{-2}$  at  $20^\circ\text{C}$ ), and the particle diameter ( $d_p$ ):  $\tau \approx 9.3 \times 10^6 d_p^2$ . The characteristic time scale for the wind field can be estimated to be of the order of  $t \approx 10^{-1} \text{ s}$ , which, in combination with an aerosol diameter  $d_p < 2 \times 10^{-5} \text{ m}$ , gives  $St \approx 3.7 \times 10^{-2} \ll 1$ .

Laser Doppler anemometry (LDA) was used to measure the fluctuating velocity field whereas the concentration measurements were conducted using a fast flame ionisation



**Fig. 2** An overview of the measuring locations. The *points* labeled  $p_{NN}$  represent vertical measuring lines for the velocity, and the *lines* labeled  $c_{NN}$  represent measuring lines for the concentration field. The global wind direction is along the  $y$ -axis

366 detector (FFID) at a sampling frequency of 200 Hz. Ground level concentrations, 0.010 m  
 367 above the model surface (corresponding to approximately 3 m above the ground in full scale),  
 368 were measured, as well as vertical variations of the concentration. The data were collected in  
 369 blocks, in total 200 at each measurement position, in order to ensure statistical convergence  
 370 of the data.

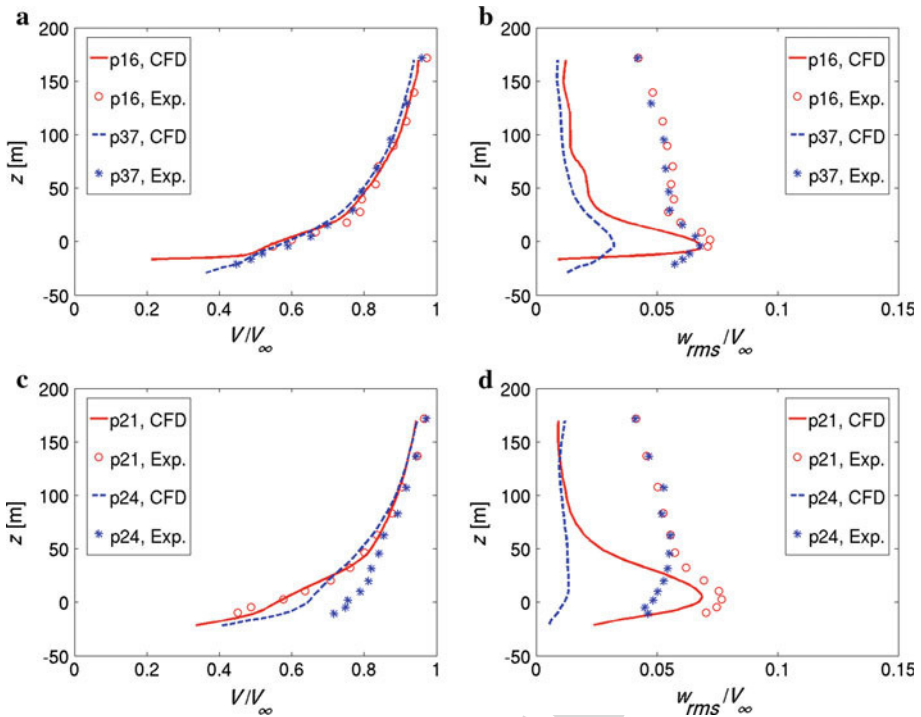
## 371 5 Results and Discussion

372 Considering the complexity and dimensions of the geometry, the results generally agreed  
 373 well with experimental data, as will be seen shortly. The greatest discrepancies, seen in the  
 374 turbulent fluctuations of the flow, are most likely due to somewhat inaccurate boundary con-  
 375 ditions. The measuring locations used in this study are indicated in Fig. 2. All distances in  
 376 all directions in the following figures are measured from the centre of the surface of pond 1  
 377 in Fig. 2.

378 The flow field results are scaled with the inlet free-stream velocity, whereas the contami-  
 379 nant field results are scaled with the mean concentration source strength. However, the source  
 380 characteristics of the simulations and the experiments differ (cf. Sects. 4 and 3.4.1). There-  
 381 fore, an additional scaling is used, in order to scale the experimental concentration field to the  
 382 simulation field. This scaling factor is applied globally and is based on the average ratio of  
 383 experiment to simulation plume maximum values for six spanwise lines ( $c_{03}$ ,  $c_{05}$  and  $c_{07}$   
 384 and three other similar lines spread uniformly downstream the source). The scaling factor is  
 385 set as  $2.8057 \times 10^{-4}$ .

### 386 5.1 Flow Field

387 Figure 3a and c shows mean streamwise velocities along the vertical lines situated at  $p_{03}$   
 388 and  $p_{16}$ , and  $p_{21}$  and  $p_{24}$ , respectively. Note that in the former figure, the two locations  
 389 lie along a streamwise line, and in the latter figure, the locations lie along a spanwise line



**Fig. 3** Mean streamwise and root-mean-squared vertical velocities, scaled with mean inlet free-stream velocity. Vertical distance measured from the surface of pond 1, cf. Fig. 2

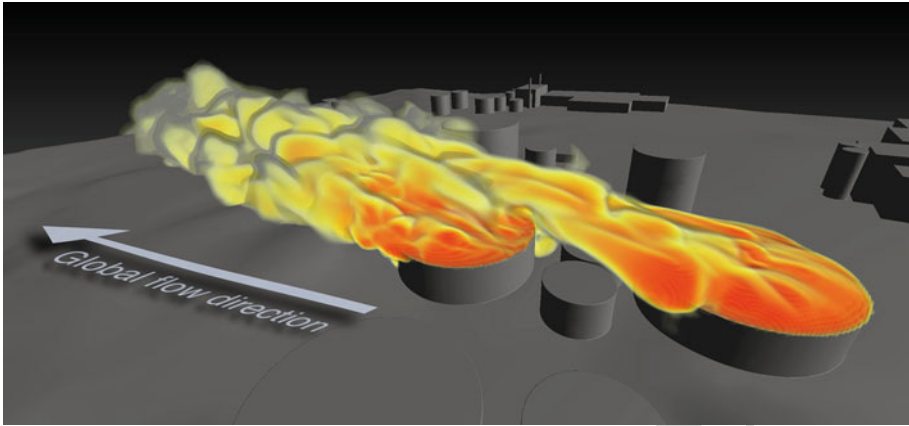
(cf. Fig. 2). The simulation results fit well with the experimental data, although the speed along p24 is somewhat under-predicted. Considering its location in the terrain with very few upstream building structures, this discrepancy is likely to be due to errors in the inlet conditions.

The root-mean-squared (r.m.s.) values of vertical velocity are shown in Fig. 3b (locations p03 and p16) and 3d (locations p21 and p24). The general trends agree, and there is a consistency between the different locations, but the actual values from the simulation are generally much lower than those from the wind-tunnel experiment, particularly far from the ground ( $z > 50$ ).

Interestingly, the r.m.s. levels of vertical velocity show a better agreement closer to the ground, i.e. in the more complex geometry, where the inlet conditions have less effect. This indicates that the turbulence levels in the simulation are considerably under-predicted due to errors in the boundary conditions rather than predicted dynamic evolution per se. This affects the solution throughout the domain, but to a much lesser degree at locations where turbulence is generated by geometry. For concentration predictions, the r.m.s. velocities at about the plume height (around  $z = 0$  in the figures) are most significant. Here, the values agree well, at least in areas containing complex structures (such as p16 and p21).

## 5.2 Contaminant Field

For illustrational purposes, a snapshot of the instantaneous concentration plume is shown in Fig. 4. Turbulent concentration fluctuations are evident in the figure.



**Fig. 4** A snapshot of the instantaneous concentration plume. Red colour high concentration; yellow colour low concentration. Created with VoluViz (Gaarder and Helgeland 2002)

410 Mean concentrations are shown in Fig. 5. Figure 5a shows the spanwise measuring lines  
 411  $c03$ ,  $c05$  and  $c07$ , Fig. 5c shows the vertical measuring lines  $c09$ ,  $c11$  and  $c13$ , and Fig.  
 412 5e shows the vertical measuring lines  $c15$ ,  $c16$  and  $c17$ , cf. Fig. 2. The spanwise lines are  
 413 always located 3 m above the ground. As can be seen, the simulation results agree well with  
 414 the experimental data, especially  $c05$ ,  $c07$ ,  $c11$  and  $c13$ .

415 In Fig. 5a, the plume width is somewhat under-predicted in  $c03$ , but not significantly.  
 416 Otherwise, the plume width is correctly computed. The ratios of simulated to measured plume  
 417 width are 0.86 for  $c03$ , 0.90 for  $c05$  and 0.93 for  $c07$ , measured at a height corresponding to  
 418 half of where the concentration peaks in each case, using linear interpolation of experimental  
 419 data. The decay of the peak concentration fits well, but the location of the peak is somewhat  
 420 misaligned.

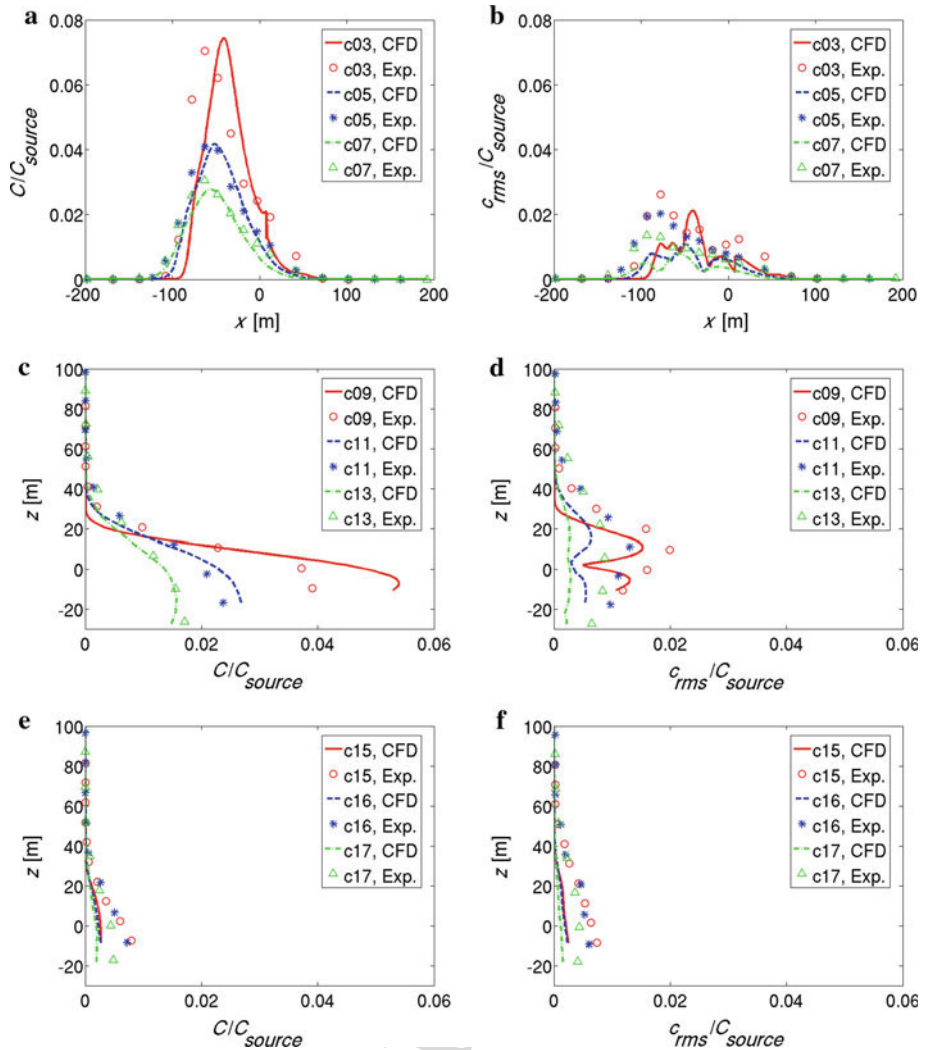
421 The concentration along  $c09$  is over-predicted close to the ground, cf. Fig. 5c, possi-  
 422 bly due to the plume misalignment discussed above. In contrast, the values higher up are  
 423 somewhat under-predicted, which might be caused by the under-prediction of  $w_{rms}$ . cf. Fig.  
 424 3b.

425 The mean concentrations in Fig. 5e are consistent with the wind-tunnel measurements, but  
 426 they are slightly under-predicted. Most likely this can be attributed to the slight misalignment  
 427 of the concentration plume.

428 Figure 5b, d and f displays rms concentrations along the same vertical lines as the pre-  
 429 vious mean concentration plots. Figure 5b shows the spanwise measuring lines  $c03$ ,  $c05$   
 430 and  $c07$ , Fig. 5d shows the vertical measuring lines  $c09$ ,  $c11$  and  $c13$ , and Fig. 5f shows  
 431 the vertical measuring lines  $c15$ ,  $c16$  and  $c17$ . The predictions are consistent with each  
 432 other, and the effect of the slight plume misalignment is apparent here as well.  $c_{rms}/C$  seems  
 433 to be under-predicted somewhat throughout the domain compared to the experiments. On  
 434 average, the CFD r.m.s. concentration is 42% of the mean for the listed locations, whereas  
 435 in the experiments the r.m.s. concentration is 52% of the mean (both cases measured where  
 436 the concentration is at its maximum). The difference between numerical and experimental  
 437 r.m.s. to mean ratios is considered relatively minor.

438 Finally, Fig. 6 shows the decay of the peak concentration of the plume as it propagates  
 439 through the domain. More specifically, the maximum values along several spanwise lines  
 440 (including  $c03$ ,  $c05$  and  $c07$ ) are plotted versus their location in the streamwise direction.





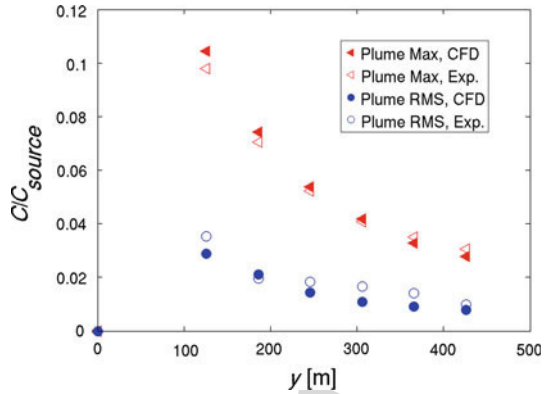
**Fig. 5** Mean and root-mean-squared concentration values for various measuring lines, cf. Fig. 2. Distances (both vertical and horizontal) are measured from the surface centre of pond 1, also cf. Fig. 2

441 Note, (i) that this plot thus corrects for the plume misalignment, by following the plume maximum for both the experimental and computational case, and (ii) that the maximum value of  $C$  is sought out only in the spanwise direction, not vertically. As mentioned previously, the spanwise lines lie 3 m above the ground. The most interesting feature of the results in Fig. 6  
 442  
 443  
 444  
 445 is how well the decay is predicted in the computational model.

#### 446 5.2.1 Statistical Analysis

447 In this section the overall performance of the simulations will be considered by using statistical performance measures to compare the simulated and experimental concentrations of the tracer gas, which was represented in the simulations by a passive scalar.  
 448  
 449

**Fig. 6** Maximum values along several spanwise lines (including  $c03$ ,  $c05$  and  $c07$ , cf. Fig. 2) are plotted versus their location in the streamwise direction. Distance is measured from pond 1, cf. Fig. 2. The second, fourth and sixth locations along the  $y$ -axis correspond to the locations of the spanwise lines  $c03$ ,  $c05$  and  $c07$



A large number of statistical performance measures have been proposed and used to evaluate dispersion models, see for example Chang and Hanna (2004). In the present case, the comparison will be performed using the metrics factor of two ( $Fa2$ ), mean relative bias ( $MRB$ ), fractional bias ( $FB$ ), mean relative square error ( $MRSE$ ), and normalized mean square error ( $NMSE$ ), in line with the recommendations from the SMEDIS project (Carissimo et al. 2001) and the more recent COST 732 action (Schatzmann et al. 2010; Di Sabatino et al. 2011).

In the following discussion it is assumed that one has a set of predictions and observations, denoted by  $C'_p$  and  $C'_o$ , respectively. Measurement positions that lie close to the edge of the plume can make a disproportionally large impact on the computed values of the statistical performance measures. To overcome this, a threshold value is set, and the data are processed such that all values that are smaller than the threshold are set to the threshold value. Mathematically, one can write

$$C = \max(C', C_{\text{thresh}}), \tag{8}$$

where  $C$  is the processed value that will be used in the comparison. In the present work the threshold  $C_{\text{thresh}} = 10^{-4}$  is employed. This is consistent with the sensitivity of the mean concentration measurements in the wind tunnel.

The results of dispersion models have traditionally been considered adequate if they give results within a  $Fa2$  of the reference data, i.e. if one has

$$C_o/2 \leq C_p \leq 2C_o. \tag{9}$$

The  $Fa2$  performance measure is defined as the fraction of monitor locations where the following holds:

$$Fa2 = \frac{N|_{C_o/2 \leq C_p \leq 2C_o}}{N_{\text{tot}}}. \tag{10}$$

For the present simulation  $Fa2 = 0.757$  is obtained, i.e. the predicted results are in the factor-of-two band for 76% of the measurement positions.

The  $Fa2$  metric only gives the fraction of the measurements that lie within the factor-of-two band. To obtain more information, other metrics must be considered. The  $MRB$  and  $FB$  are measures of over- or under-prediction defined by

Author Proof

$$MRB = \left\langle \frac{2(C_p - C_o)}{C_p + C_o} \right\rangle, \quad (11)$$

$$FB = \frac{2\langle C_o - C_p \rangle}{\langle C_o + C_p \rangle}, \quad (12)$$

where the angular brackets denote an average over the entire dataset. Note the difference in sign of the two metrics, such that under-prediction is represented by a negative  $MRB$  value and a positive  $FB$  value and vice versa for over-prediction.

Related metrics are the  $MRSE$  and the  $NMSE$ . In additions to systematic errors, these metrics also quantify the variance of the difference between the compared datasets and is thus a measure of the scatter in the comparison of the predicted and observed values. The  $MRSE$  and  $NMSE$  are defined by

$$MRSE = 4 \left\langle \left( \frac{C_p - C_o}{C_p + C_o} \right)^2 \right\rangle, \quad (13)$$

$$NMSE = \frac{\langle (C_o - C_p)^2 \rangle}{\langle C_o \rangle \langle C_p \rangle}. \quad (14)$$

For the present simulation,  $MRB = -0.284$ ,  $FB = 0.032$  and  $MRSE = 0.385$ ,  $NMSE = 3.018$  are obtained. We note in passing that the computed  $MRB$  value corresponds to a systematic under-prediction of 25%, whereas over- and under-prediction almost cancel out for the  $FB$  metric.

When considered together, the  $MRB/MRSE$  and  $FB/NMSE$  pairs give information both on the level of over- and under-prediction and to what extent the predictions are consistent with the observations. This can be shown by noting that always,

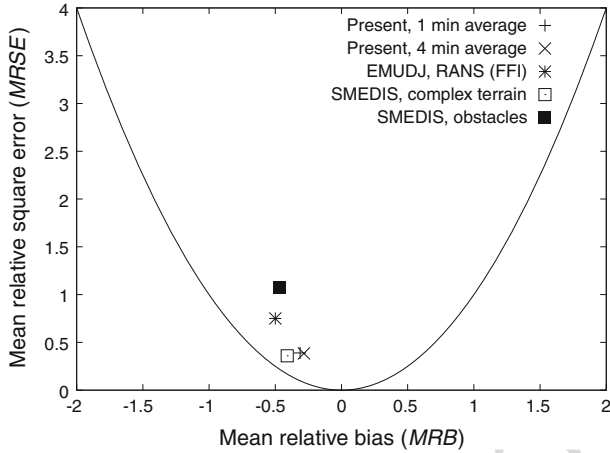
$$MRSE \geq MRB^2, \quad (15)$$

$$NMSE \geq \frac{4FB^2}{4 - FB^2}. \quad (16)$$

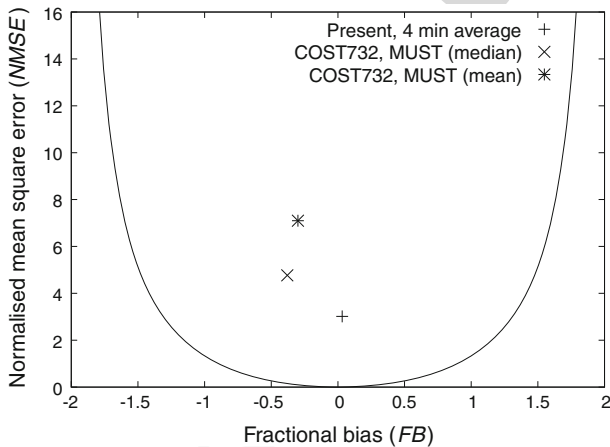
with equality in the case of systematic over- or under-prediction. If  $MRB$  versus  $MRSE$  or  $FB$  versus  $NMSE$  is plotted, there will be an 'ideal' trend curve given by the lower bounds given above. The extent that a model result is consistent with the reference data can then be assessed by the distance to the trend curve.

The SMEDIS project (Carissimo et al. 2001) comprised a large number of cases and reference datasets. Of these, in particular, the cases EMUDJ and EMUNJ are of a comparable complexity to the present study. A scaled-down model of an industrial installation—located in Amlwch, Wales—and the surrounding countryside was constructed and wind-tunnel experiments were performed in the Surrey EnFlo facility during the Evaluation of Model Uncertainty (EMU) project (Cowan and Robbins 1996). The experimental data were later post-processed and scaled back to real scale by the SMEDIS project.

In Fig. 7 the computed  $MRB$  versus  $MRSE$  for the present calculation is shown. For comparison we also include results from a recent simulation of the EMUDJ case (Gjesdal et al. 2008) and average results from simulations performed by the SMEDIS project for the scenario categories *complex terrain* and *obstacles* (Carissimo et al. 2001). Note that the majority of *terrain* scenarios in SMEDIS were not from obstructed geometries. As a consequence, the performance measures for these cases may possibly be skewed away from the EMU scenarios.



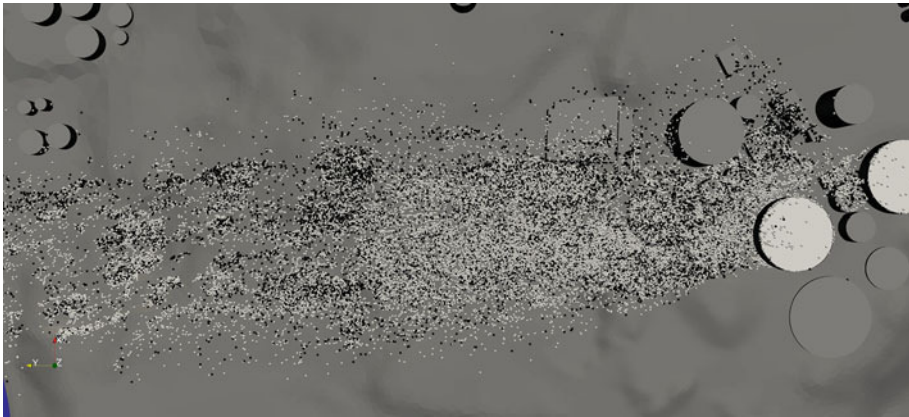
**Fig. 7** Statistical performance measures, *MRB* versus *MRSE*, for the present simulation compared to an earlier simulation of the EMUDJ reference cases and SMEDIS results for scenarios characterized by complex terrain and obstacles, respectively



**Fig. 8** Statistical performance measures, *FB* versus *NMSE*, for the present simulation compared to the corresponding median and mean quantities from the COST 732 comparison exercise of dispersion in the MUST obstacle array geometry

516 It is therefore probable that the performance measures for the EMU cases in SMEDIS lie  
 517 somewhere between the two points.

518 Recently, the COST 732 Action (Schatzmann et al. 2010; Di Sabatino et al. 2011) per-  
 519 formed a comparison of 18 different simulations of a dispersion scenario in the MUST  
 520 obstacle array geometry. In Fig. 8 the computed *FB* and *NMSE* from the present simulations  
 521 and the median and mean of the corresponding quantities from the COST 732 exercise are  
 522 shown.



**Fig. 9** Accumulated deposition at the end of the discrete particle simulation. *Dark colour* heavy particles; *light colour* light particles

### 523 5.3 Ground Deposition

524 Discrete particle transport was simulated primarily in order to predict particle deposition.  
 525 More specifically, it is of interest to see if both the large ( $18\ \mu\text{m}$ ) and small ( $2\ \mu\text{m}$ ) particles  
 526 are deposited at approximately the same rate, or if only the smaller aerosols are transported  
 527 over longer distances. No experimental reference data are available for particle deposition,  
 528 but in light of the previous results, it is likely that the simulation provides a good indication  
 529 as to how particles emitted from the aeration ponds are deposited in the domain. Note that  
 530 the following results are only tentative, as a more complete analysis requires knowledge of  
 531 second-order effects (evaporation, agglomeration etc.), a much higher number of particles,  
 532 and a longer release time, in order to approach statistical convergence.

533 Figure 9 shows the deposition on the ground and building roofs after all (>99%) released  
 534 particles have left the domain or been deposited. Clearly, the deposition is chaotic, indicating  
 535 the presence of turbulence. Within the near-field domain simulated here, there seems to be  
 536 no significant difference in the deposition patterns for large and small particles. The simu-  
 537 lation results show that 23% of the released particles are deposited, of which 58% are large  
 538 particles. Hence, there is no significant differences between the deposition of large and small  
 539 particles within the first few hundred metres downstream of the ponds.

### 540 6 Concluding Remarks

541 It has been demonstrated that computational fluid dynamics (CFD) can faithfully predict the  
 542 aerial dispersion of aerosols at low Stokes numbers in complex industrial/urban terrain in  
 543 the case of neutral conditions and flow in the high Reynolds number range. There is gener-  
 544 ally a good correspondence between the simulations and the experimental measurements,  
 545 especially in the near-surface layer (0–50 m from the ground). In particular, the predicted  
 546 decay of both the mean and r.m.s. concentrations are in excellent agreement with wind-tun-  
 547 nel experiments; the mean concentration 200 m downwind the plume is approximately 2% of  
 548 the source strength 3 m from the ground. The r.m.s. concentration levels play an increasingly  
 549 important role downwind as the mean concentration is reduced.

The results show that the plume emitted from the ponds remain narrow up to at least 500 m downwind (100–150 m wide), and that it grows very slowly in the vertical direction (40–50 m in height at 400 m downwind the ponds). Also, the maximum predicted/measured mean concentration remain very close to the ground ( $\leq 3$  m), whereas the concentration fluctuations attain their maxima 10–20 m above the ground. Vertical r.m.s. velocity fluctuations peak at around 10–30 m above the ground as well, close to where the r.m.s. concentration peaks. Finally, there seemed to be little difference in the deposition rates of large and small particles in the near-field. The results of this study also indicates that the rate at which aerosols emitted from the ponds are deposited constitutes a significant effect with respect to the expected overall concentration level further downwind.

In summary, CFD has proven to be a useful tool for analyzing, predicting and understanding aerosol dispersion in the neutral atmospheric boundary layer. The present study also demonstrates the large impact of the chosen inflow conditions in areas of the domain with little complex geometry (i.e. at high elevations). Further extensions of the methodology used herein, e.g. to stratified boundary layers will be investigated in a future study, in which stratified inflow will have to be generated in a separate periodic domain with ground heating or cooling.

**Acknowledgments** Part of this work was financially supported by Borregaard Ind. Ltd. and by the Norwegian Research Council through the Norwegian Center of Excellence Center for Biomedical Computing. We would like to thank Drs Anders Helgeland (FFI), Øyvind Andreassen (FFI), Janet M. Blatny (FFI), and Viggo Waagen (Borregaard Ind. Ltd.). The staff at the EnFlo laboratory at the University of Surrey, Guildford, UK, Prof Alan Robins and Dr Paul Hayden, are also acknowledged for their support during the project. Finally, the authors are grateful for valuable feedback from the reviewers.

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